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1 **Influence of the Leeuwin Current on the epipelagic euphausiid assemblages of the south-east**  
2 **Indian Ocean**

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8 **Abstract**

9 The Leeuwin Current is an anomalous eastern boundary current which transports warm, low salinity water poleward off  
10 Western Australia. This study investigated epi-pelagic euphausiids in the Leeuwin Current system from 22°S - 34°S  
11 focusing on the latitudinal gradient in species richness and whether variability in euphausiid assemblages was associated  
12 with an increase in seawater density across latitude. Twenty-eight species of euphausiids (including five new records) were  
13 identified from the study area. Species richness remained relatively constant across latitude as the distribution of seven  
14 tropical species, including the dominant *Pseudeuphausia latifrons*, extended south of 30°S. Euphausiid assemblages from  
15 the northern shelf stations were distinct from the oceanic, shelf break and southern shelf assemblages which, though  
16 clustering together, showed evidence of latitudinal shifts and day/night influence, particularly at oceanic  
17 stations. Distance-based linear models confirmed that, of the environmental variables examined, mean seawater density  
18 was the only significant explanatory variable accounting for 32%, 27% and 71% of the variation for shelf, shelf break and  
19 night oceanic assemblages, respectively. This study provides the first account of the diversity, distribution and abundance  
20 of euphausiids along the entire western seaboard of Australia and enhances understanding of the influence of the Leeuwin  
21 Current on holoplanktonic biota.

22 **KEY WORDS:** Oceanography, boundary current, Euphausiacea, water mass, krill

23 **Introduction**

24 Euphausiids are holoplanktonic, vertically migrating crustaceans and the distributions of the 86 species known from the  
25 world's oceans are relatively well documented (Brinton et al., 2000). They generally have affinities with tropical,  
26 subtropical, temperate or polar marine environments, although some are cosmopolitan in their distribution. Euphausiids  
27 follow the same global trend as many other marine biota, with the number of species decreasing with increasing latitude  
28 (Letessier et al., 2009; Tittensor et al., 2010; Letessier et al., 2011). However, their distribution patterns are known to be

29 influenced by ocean currents (Mauchline & Fisher, 1969; Pillar et al., 1989; Barange & Pillar, 1992) and, in particular,  
30 western boundary currents can disrupt the latitudinal gradient in species richness by transporting waters with tropical  
31 characteristics and complement of plankton towards temperate environments (Wang & Chen, 1963; Griffiths, 1979;  
32 Gibbons et al., 1995). In the south-east Indian Ocean, the Leeuwin Current (LC), an anomalous, poleward flowing eastern  
33 boundary current, occurs off Western Australia. This current enables tropical corals, fishes and other taxa to exist as far  
34 south as the Houtman Abrolhos Islands (29°S) and Rottnest Island (32°S) (Maxwell & Cresswell, 1981; Wilson & Allen,  
35 1987; Hutchins & Pearce, 1994; Fox & Beckley, 2005) but its influence has not been investigated with respect to the  
36 distribution of epi-pelagic euphausiids.

37 The LC comprises warm, low salinity, Tropical Surface Water (TSW) sourced from the Indo-Australian basin, with inputs  
38 from the Indonesian Throughflow, the South Java Current, the South Equatorial Current and the Eastern Gyral Current  
39 (Meyers et al., 1995; Feng & Wijffels, 2002; Domingues et al., 2007). As the LC flows poleward along the shelf break,  
40 cooler, more saline Sub Tropical Surface Water (STSW) is entrained into the current and, together with evaporative  
41 cooling, results in longshore temperature and salinity gradients from north to south along the Western Australian coast  
42 (Woo et al., 2006; Weller et al., 2011). The strength of the LC is intensified during the austral autumn/winter months,  
43 when maximum poleward geostrophic transport can reach 5 Sv (Feng et al., 2003), and the current, after passing Cape  
44 Leeuwin (34°S), extends eastward across the Great Australian Bight towards Tasmania (Cresswell & Golding, 1980;  
45 Ridgway & Condie, 2004).

46 Studies on the euphausiids within the LC are few, particularly for a region that supports southern blue fin tuna, migratory  
47 baleen whale populations and globally important colonies of seabirds, all of which are known consumers of euphausiids  
48 (Surman & Wooller, 2003; Rennie et al., 2009; Itoh et al., 2011). Four euphausiid species, namely, *Pseudeuphausia*  
49 *latifrons* (Sars, 1883), *Euphausia hemigibba* Hansen, 1910, *Euphausia recurva* Hansen, 1905 and *Thysanopoda*  
50 *tricuspidata* Milne-Edwards, 1837, have been previously reported from the LC (Wilson et al., 2003; Rennie et al., 2009),  
51 and 22 species were identified within a mesoscale eddy of the LC located between 31°S - 34°S and 112°E - 115°E (Sutton  
52 et al., 2015). Eleven species have been recorded in the headwaters of the LC between 20°S - 26°S, 105°E - 118°E  
53 (Taniguchi, 1974) and, in oceanic waters west of the LC, 31 euphausiid species were reported to occur along the 110°E  
54 meridian between 9°S - 32°S during the 1962 - 1965 International Indian Ocean Expedition (McWilliam, 1977).

55 The primary aim of this study was to examine the diversity, distribution and abundance of epipelagic euphausiids off the  
56 western seaboard of the Australian continent (22°S - 34°S) and relate these to prevailing oceanographic conditions. As the  
57 study area is dominated by the LC, the specific research questions were: a) does species richness of epi-pelagic

58 euphausiids decline with increasing latitude and b) are epi-pelagic euphausiid assemblages off Western Australia  
59 structured by the change in seawater density across latitude?

## 60 **Methods**

### 61 **Study area**

62 A multi-disciplinary survey of the LC system between 22°S - 34°S and 111°E - 115°E was conducted during the late  
63 austral autumn from 16 May - 5 June 2007 aboard the R.V. *Southern Surveyor* (voyage SS04/07). Thirteen cross-shelf  
64 transects were sampled at each degree of latitude off the coast of Western Australia, encompassing shelf (~ 50 m depth),  
65 shelf break (~ 200 m) and oceanic (1000 - 2000 m) stations (Fig. 1). The concurrent studies of physical oceanography  
66 (Weller et al., 2011), nutrients (Thompson et al., 2011), primary production (Lourey et al., 2012), larval fishes (Holliday et  
67 al., 2012) and chaetognaths (Buchanan & Beckley 2016) identified physical and biological structuring in a system with  
68 dynamic longshore and cross-shelf transport.

### 69 **Oceanographic and biological sampling**

70 Abiotic and biotic properties of the water column at each station were sampled using a Seabird SBE 911 instrument (CTD)  
71 mounted in a 24 Niskin bottle rosette. The CTD was equipped with a dual temperature and conductivity sensor, an oxygen  
72 sensor and a Chelsea TGI fluorometer. Chlorophyll *a* was determined from water samples collected at the surface, 10, 25,  
73 50, 75 and 100 m depths, and at the chlorophyll *a* maximum when present (Lourey et al., 2012). Temperature and salinity  
74 were used to calculate a mean seawater density for each station (0 – 150 m depth), expressed at sigma-t ( $\sigma_t$ ).

75 After oceanographic measurements were taken, irrespective of whether it was day or night, zooplankton samples were  
76 collected using a bongo net (355  $\mu$ m and 100  $\mu$ m meshes, net diameters of 0.5 m). Two replicate oblique tows from a  
77 maximum depth of 150 m (or shallower on the shelf) to the surface were conducted at each station. The bongo net was  
78 towed at a speed of 2 knots for approximately 15 minutes. To determine the volume of water sampled, General Oceanics  
79 flowmeters were mounted in the mouth of both nets and were connected to an electronic interface in order to monitor the  
80 tow profile and volume of seawater sampled. Zooplankton samples were preserved in 5% buffered formaldehyde in  
81 seawater and the 355  $\mu$ m mesh samples were used for identification and quantification of euphausiids. As the plankton  
82 samples were required to be intact for other projects, estimates of mesozooplankton abundance were made by pouring the  
83 zooplankton sample through a 1 mm sieve and measuring the settled volume of the < 1 mm zooplankton in a graduated  
84 cylinder after a 24 h period (Gibbons, 1999; Suthers & Rissik, 2009); settled volumes were expressed in mL m<sup>-3</sup>.

## 85 **Euphausiid identification**

86 Immature and mature euphausiids were counted using a dissecting microscope and identified using relevant literature  
87 (Baker et al., 1990, Brinton et al., 2000). Zooplankton was subsampled with a Folsom splitter and a minimum of 200  
88 euphausiids (100 immature and 100 mature specimens of all species combined) were counted from the sub-samples before  
89 estimates were made of the total concentration (Gibbons, 1999). Concentrations of euphausiids were expressed as the  
90 average number of individuals per 1000 m<sup>-3</sup> ± standard error. Damaged or indistinguishable immature and mature  
91 specimens were grouped as unidentified. Euphausiid species were classified as tropical, subtropical or temperate based  
92 upon geographical distributions given by Brinton et al. (2000).

## 93 **Statistical analyses**

94 As euphausiids undergo pronounced diel vertical migration, stations were classified *a priori* as day or night. Univariate  
95 analyses were performed on total euphausiid concentrations using the non-parametric Kruskal-Wallis test for independent  
96 samples (IBM SPSS Statistics 21), to assess if total concentrations differed across day and night and across isobath (shelf,  
97 shelf break and oceanic).

98 For multivariate analyses of euphausiid assemblages, the PRIMER v6 PERMANOVA+ software package was used  
99 (Anderson et al., 2008; Clarke & Gorley, 2015). A fourth root transformation was applied to euphausiid concentrations to  
100 reduce the relative importance of abundant species, and a Bray-Curtis resemblance matrix was constructed. Non-metric  
101 multidimensional scaling ordination was used to assess the spatial relationships among assemblages, and one-way analysis  
102 of similarity (ANOSIM) was used to test for differences in assemblages across day and night for each isobath (Clark &  
103 Warwick, 2001).

104 To determine which abiotic and biotic properties of the water column structured euphausiid assemblages, mean seawater  
105 density (0 - 150 m), mean dissolved oxygen (0 -150 m), surface chlorophyll *a*, depth integrated chlorophyll *a* (0 - 100 m)  
106 and mesozooplankton settled volume (0 - 150 m), were correlated with euphausiid assemblages using distance-based linear  
107 models (DISTLM) (Anderson, 2001). These models were constructed using the step-wise selection procedure with the  
108 adjusted R<sup>2</sup> as the selection criterion, and were used to determine the subset of environmental variables that best correlated  
109 with euphausiid assemblages. Square root transformations were applied to the environmental data, and the data were  
110 normalised and constructed into a resemblance matrix based on Euclidean distance. Draftsman plots revealed if co-  
111 correlation existed between environmental variables ( $r^2 > 0.7$ ), which resulted in the removal of mean dissolved oxygen  
112 due to its positive correlation with mean seawater density.

## 113 **Results**

### 114 **The oceanographic environment**

115 The LC was the dominant feature in the south-eastern Indian Ocean during May - June 2007 and was most evident along  
116 the shelf break where it transported warm waters southwards from 22°S to 34°S (Fig. 1). At shelf and shelf break stations  
117 the water column was generally well-mixed, but for oceanic stations as far north as 25°S, the water column was more  
118 stratified with shallower mixed layers (Fig. 2). For all stations, a gradual shift from a TSW signature ( $> 22^{\circ}\text{C}$ ,  $< 35.5$  psu,  
119  $22.9 - 24.6 \sigma_t$ ) to a STSW signature ( $< 22^{\circ}\text{C}$ ,  $> 35.7$  psu,  $24.8 - 26.2 \sigma_t$ ) was evident (Fig. 2; Fig. 3). In the top 150 m of  
120 the water column (matching the extent of epi-pelagic zooplankton sampling), a decrease in mean temperature  
121 corresponded to an increase in mean salinity and mean dissolved oxygen and overall, an increase in mean seawater density  
122 occurred with an increase in latitude (Fig. 3).

123 In general, surface chlorophyll *a* and depth-integrated chlorophyll *a* increased from north to south (Fig. 4a & b).  
124 Mesozooplankton settled volume decreased from  $2.31 \text{ ml m}^{-3}$  at 22°S to  $0.40 \text{ ml m}^{-3}$  at 34°S for shelf stations, but showed  
125 no obvious patterns for shelf break and oceanic stations (Fig. 4c).

### 126 **Euphausiid diversity, abundance and distribution**

127 A total of 28 euphausiid species was identified from the LC study area (Fig. 5a) and overall, there was no evidence of a  
128 decline in species richness with increasing latitude ( $r^2 = 0.0005$ ) (Fig. 5b). Five of these species were new records for the  
129 south east Indian Ocean, namely, *Euphausia fallax* Hansen, 1916, *Nyctiphanes australis* Sars, 1883, *Stylocheiron indicum*  
130 Silas & Mathew, 1967, *Stylocheiron insulare* Hansen, 1910, and *Stylocheiron robustum* Brinton, 1962 (Fig. 5a). Of the  
131 nine tropical species identified across the study area (*Euphausia diomedea* Ortmann, 1894, *Euphausia fallax* Hansen,  
132 1916, *Euphausia sanzoi* Torelli, 1934, *P. latifrons*, *S. indicum*, *S. insulare*, *Stylocheiron microphthalmum* Hansen, 1910,  
133 *Thysanopoda astylata* Brinton, 1975 and *T. tricuspida*), all except *S. indicum* and *T. tricuspida* were recorded south of  
134 30°S (Fig. 5). The number of species recorded was higher for oceanic (26 in total) and shelf break (22 in total) stations  
135 than shelf stations (14 in total) (Fig. 6a). The occurrence of some species was also affected by whether the samples were  
136 collected during the day or night; only 17 species were collected from day stations whereas all 28 species were recorded at  
137 night stations.

138 The highest total euphausiid concentrations were recorded at shelf stations at 22°S and 27°S, reaching  $50,858 \text{ inds} \cdot 1000$   
139  $\text{m}^{-3}$  and  $16,649 \text{ inds} \cdot 1000 \text{ m}^{-3}$ , respectively (Fig. 6b); these samples were mainly comprised of juvenile *P. latifrons*.  
140 Generally, though, total concentrations at all other stations ranged between  $130 - 4191 \text{ inds} \cdot 1000 \text{ m}^{-3}$ , and decreased from

141 north to south. More specimens were caught during the night ( $1671 \pm 251$  inds·1000 m<sup>-3</sup>) than during the day ( $820 \pm 169$   
142 inds·1000 m<sup>-3</sup>) ( $p = 0.010$ ,  $n = 39$ , Kruskal-Wallis), and there were no significant differences in total concentration across  
143 isobath ( $p = 0.676$ ,  $n = 39$ , Kruskal-Wallis).

144 *Pseudeuphausia latifrons*, *E. recurva* and *Stylocheiron carinatum* Sars, 1883 were the most common and abundant species  
145 throughout the study area and together, they accounted for over 75% of euphausiids. *Pseudeuphausia latifrons*, a tropical  
146 neritic species, was identified from every station sampled in the study area and concentrations decreased from north to  
147 south (Fig. 7a). The subtropical/temperate oceanic species, *E. recurva*, was mostly identified from stations in the southern  
148 part of the study area (29°S - 34°S), but with some occurrences between 25°S - 27°S (Fig. 7b); concentrations increased  
149 towards the south. *Stylocheiron carinatum* is typically cosmopolitan and it was collected at most stations throughout the  
150 study area with similar concentrations across latitude (Fig. 7c).

### 151 **Euphausiid assemblages and environmental correlations**

152 A nMDS ordination of euphausiid assemblages showed that shelf stations, particularly those from the north of the study  
153 area, were largely distinct from shelf break and oceanic stations which were clustered together (Fig. 8a). Whether stations  
154 were sampled during the day or night (Fig. 8b) was not significant for shelf assemblages (ANOSIM  $R = 0.074$ ,  $P = 0.219$ ,  
155  $n = 13$ ) or shelf break assemblages ( $R = 0.212$ ,  $P = 0.061$ ,  $n = 13$ ) but was significant for oceanic assemblages ( $R = 0.915$ ,  
156  $P = 0.006$ ,  $n = 13$ ).

157 In the distance-based linear model for shelf assemblages, all four environmental variables together explained 48% of the  
158 total variation (Table 1). Mean seawater density was, however, the only significant explanatory variable accounting for  
159 32% ( $P = 0.002$ ). For the fitted model, the distance-based redundancy bi-plot showed axes 1 and 2 to explain 69% and  
160 24% of the variation, respectively (Fig. 9a). The change in euphausiid assemblages from 22°S - 34°S corresponded with  
161 increasing seawater density.

162 For euphausiid assemblages along the shelf break, all four environmental variables together explained 61% of the total  
163 variation (Table 1). Mean seawater density was again the only significant variable accounting for 27% ( $P = 0.001$ ). For the  
164 fitted model, axes 1 and 2 explained 53% and 26% of the variation, respectively (Fig. 9b). Similarly, the change in  
165 euphausiid assemblages across latitude corresponded with increasing seawater density.

166 The significant day/night effect on oceanic euphausiid assemblages was taken into account for the distance-based linear  
167 model. For assemblages sampled during the day, although the four environmental variables explained 86% of the  
168 variation, none of the variables was significant (Table 1). For assemblages sampled at night, the four environmental

169 variables explained 77% of the variation (Table 1). Mean seawater density was the only significant explanatory variable  
170 explaining 51% of the variation ( $P = 0.004$ ). For the fitted model, axes 1 and 2 explained 68% and 20% of the variation in  
171 night assemblages, respectively (Fig. 9c).

## 172 **Discussion**

173 This study ascertained that, over the 22° - 34°S study area off Western Australia, the total number of species of epi-pelagic  
174 euphausiids did not decline with increasing latitude but remained relatively stable confirming that the anomalous LC  
175 disrupts the expected latitudinal gradient in euphausiid species richness. However, the mix and concentrations of species  
176 contributing to euphausiid assemblages at shelf, shelf break and oceanic stations over the study area did vary, and  
177 assemblages were shown to be significantly structured by seawater density, which increased across latitude.

178 Twenty eight euphausiid species were identified from the LC study area and comprised nine tropical, 15  
179 tropical/subtropical, three subtropical/temperate and one temperate species (Brinton et al., 2000). This mix of euphausiid  
180 species in the LC study area could be explained by entrainment of species from source waters or surrounding water  
181 masses. For example, tropical Indo-Australian endemics such as *S. insulare*, *E. fallax* and *E. sanzoi* could have been  
182 funnelled into the LC via the catchment waters of the Indonesian Throughflow. *Stylocheiron robustum*, which has a  
183 distribution across the central Indian Ocean basin (Brinton et al., 2000), was recorded at oceanic stations 25°S and 34°S,  
184 and is probably indicative of eastward flows into the LC (Domingues et al., 2007; Menezes et al., 2014). The  
185 subtropical/temperate species, *E. recurva*, *Euphausia similis* Sars, 1883, *N. australis*, and *Thysanoessa gregaria* Sars,  
186 1883 (Brinton et al., 2000), were all present in the south where the influence of STSW created a cooler, more saline  
187 environment (Domingues et al., 2007; Woo & Pattiaratchi, 2008).

188 The occurrence of the tropical species, *E. diomedea*, *E. sanzoi*, *E. fallax*, *S. insulare*, *S. microphthalma*, *T. astylata* and,  
189 particularly, *P. latifrons*, at southern latitudes confirms the influence of the southward flowing LC on the distribution of  
190 holoplanktonic biota. Such poleward dispersal of tropical euphausiid species is not unusual in western boundary currents.  
191 For example, *E. diomedea* and *P. latifrons* have been recorded to be transported northwards in the western Pacific by the  
192 Kuroshio Current (Wang & Chen, 1963, Brinton et al., 2000) and the southward dispersal of *P. latifrons* is known for both  
193 the Agulhas and East Australian Currents (Griffiths 1979; Gibbons et al. 1995).

194 In the northern hemisphere, *Nyctiphanes simplex* Hansen, 1911 has been used as a biological tracer to determine the  
195 direction of transport of warm water off Oregon during an El Niño event (Keister et al., 2005). It is proposed here that *P.*  
196 *latifrons*, a typically tropical and neritic species (Brinton et al., 2000), could be similarly used as a tracer of longshore and



197 cross-shelf transport of LC waters. Not only was *P. latifrons* found at every degree of latitude in the survey area but also at  
198 every station in shelf break and oceanic waters, which for a coastal species, could be achieved by cross-shelf transport via  
199 meanders and eddies (Holliday et al., 2011; Holliday et al., 2012; Paterson et al., 2013).

200 Numerically, three species, *P. latifrons*, *E. recurva* and *S. carinatum* were dominant and constituted 75% of all  
201 euphausiids sampled in the LC study area. It appears to be a common finding that euphausiid assemblages are numerically  
202 dominated by only a few euphausiid species, with the remaining species found in relatively low abundances. Off Baja  
203 California, *Nyctiphanes simplex* dominates assemblages (Gomez-Gutierrez et al., 1995), *Euphausia lucens* Hansen, 1905,  
204 *Euphausia hanseni* Zimmer, 1915 and *Nyctiphanes capensis* Hansen, 1911 dominate in the Benguela Current (Pillar et al.,  
205 1992), *Euphausia mucronata* Sars, 1883 dominates in the Humboldt Current system (Antezana, 2010) and *Euphausia*  
206 *pacifica* Hansen, 1911 is very abundant in the Oyashio and Kuroshio Currents (Taki, 1998; Taki, 2008).

207 Examination of the relationships between euphausiid assemblages across the study area indicated that those at northern  
208 shelf stations were distinct from the oceanic, shelf break and southern shelf assemblages. Within these nMDS groupings  
209 there was evidence of a latitudinal shift from 22° - 34°S. However, for oceanic assemblages sampled during the day, this  
210 pattern was disrupted, presumably due to the diel vertical migration of many euphausiid species over the deeper water  
211 column (Mauchline & Fisher, 1969). Assemblages sampled at the oceanic stations during the day had fewer species and  
212 lower concentrations and thus did not cluster with oceanic stations sampled at night. Depth-stratified sampling during both  
213 day and night would be required to confirm the effects of diel vertical migration on assemblages. Further, euphausiids are  
214 also known to exhibit net avoidance particularly during the day (Brinton, 1967; Wiebe et al., 1982), so use of bongo nets in  
215 this study might have influenced capture rates.

216 Euphausiids are well-known to be linked with the physical properties of water masses (Brinton, 1975; Dadon & Boltovsky,  
217 1982; Gibbons et al., 1995; Tarling et al., 1995). Seawater density is indicative of water mass and, in the LC study area,  
218 both TSW (22.9 - 24.6  $\sigma_t$ ) and STSW (24.8 - 26.2  $\sigma_t$ ) were the dominant water masses. Distance-based linear models  
219 confirmed that mean seawater density was the only significant explanatory variable accounting for 32%, 27% and 71% of  
220 the variation for shelf, shelf break and night oceanic assemblages, respectively. Lower density waters, i.e. TSW, had  
221 higher total concentrations of euphausiids, particularly of tropical species, such as *P. latifrons*, whereas higher density  
222 waters, i.e. STSW, had lower overall concentrations, but greater concentrations of subtropical and temperate species, such  
223 as *E. recurva*. Likewise, in the California Current, changes in the euphausiid assemblages and dominant species were, in  
224 part, attributed to the mixing of species from the different water masses in the region (Youngbluth, 1976).

225 Besides the physical properties of the water column, other factors, particularly food availability, are known to affect  
226 euphausiids (Youngbluth, 1975, Hu, 1978, Gibbons, 1993, Taki, 2008). As they feed on a range of phytoplankton and  
227 zooplankton species (Mauchline & Fisher, 1969), chlorophyll *a* and mesozooplankton settled volume were included as  
228 environmental variables in the distance-based linear modelling in this study. Although increases in surface and depth  
229 integrated chlorophyll *a*, and a decrease in mesozooplankton settled volume from north to south explained some of the  
230 variation in euphausiid assemblages, these were not found to be significant variables in the model.

231 Although euphausiid diversity, distribution and abundance are relatively well known across the oceans (e.g. Brinton, 2000;  
232 Letessier et al., 2009; Letessier et al., 2011), until this study, the influence of the unusual LC on euphausiids in the south-  
233 east Indian Ocean was unknown. The presence of the poleward flowing LC results in a disruption of the typical latitudinal  
234 gradient in species richness by dispersing tropical species beyond their normal distribution limits to mix with subtropical  
235 and temperate species off Western Australia. The oceanographic property of mean seawater density was established to be  
236 the best explanatory variable in structuring epi-pelagic euphausiid assemblages and emphasizes the influence of boundary  
237 currents on holoplanktonic biota.

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368 **Table 1** The percentage of the variance in euphausiid assemblages explained by each environmental explanatory variable  
369 included in the distance-based linear models for shelf, shelf break and oceanic stations. Due to the significant influence of  
370 day and night on oceanic assemblages, models were performed separately. Environmental variables explaining a  
371 significant percentage of variation are denoted by an asterisk

372 **Fig. 1** a) Sea surface temperature for the 5 June 2007, showing the warm, southward flowing Leeuwin Current (LC),  
373 Tropical Surface Water (TSW) and cooler Sub Tropical Surface Water (STSW) in the south-east Indian Ocean (adapted  
374 from Weller et al., 2011). b) Zooplankton sampling locations at shelf, shelf break and oceanic stations. The 200 m isobath  
375 is indicated by the dashed line

376 **Fig. 2** Temperature and salinity profiles (0 - 150 m) of shelf, shelf break and oceanic water columns at 22°S, 29°S and  
377 34°S in the south-east Indian Ocean. Solid line indicates temperature (°C) and dashed line indicates salinity (psu)

378 **Fig. 3** (a) Temperature and salinity profile of the water column, to 2000 m depth, across the Leeuwin Current study area in  
379 the south-east Indian Ocean. Box indicates 0 - 150 m of the water column which is expanded to show the correlations  
380 between b) mean temperature and mean salinity c) mean temperature and mean dissolved oxygen and d) latitude and mean  
381 seawater density for shelf, shelf break and oceanic environments.

382 **Fig. 4** a) Surface chlorophyll *a* concentration, b) depth integrated chlorophyll *a* (0 - 150m) and c) mesozooplankton settled  
383 volume for shelf, shelf break and oceanic stations across the Leeuwin Current study area in the south-east Indian Ocean.  
384 Grey and black circles distinguish stations sampled during the day and night, respectively

385 **Fig. 5** a) The euphausiid species and their respective latitudinal ranges (°S) recorded across the Leeuwin Current survey  
386 area in the south-east Indian Ocean in 2007. Previously known distribution ranges are given based on Brinton et al. (2000)  
387 and are indicated by \* tropical, # tropical/subtropical, ^ subtropical/temperate and + temperate. Species of a genus that  
388 could not be distinguished are indicated as sp. and spp. refers to more than one species of the same genus that could not be  
389 distinguished. b) The total number of euphausiid species identified at each latitude between 22°S - 34°S. Only euphausiids  
390 that were identified to species level were included.

391 **Fig. 6** a) The number of euphausiid species identified and b) total euphausiid concentrations for all the shelf, shelf break  
392 and oceanic stations over the Leeuwin Current study area in the south-east Indian Ocean. Grey and black circles indicate  
393 stations sampled during the day and night, respectively



394 **Fig. 7** Concentrations ( $1000 \cdot \text{m}^{-3}$ ) of the three dominant euphausiid species at shelf, shelf break and oceanic stations in the  
395 Leeuwin Current study area in the south-east Indian Ocean, a) *Pseudeuphausia latifrons*, b) *Euphausia recurva* and c)  
396 *Stylocheiron carinatum*. Scales are the same for all three species and grey and black circles indicate stations sampled  
397 during the day and night, respectively

398 **Fig. 8** nMDS ordinations showing similarities in euphausiid assemblages in the south-east Indian Ocean, defined by  
399 latitude and a) isobath (shelf, shelf break, oceanic), b) day and night

400 **Fig. 9** Distance-based redundancy bi-plots for a) shelf, b) shelf break and c) oceanic night euphausiid assemblages,  
401 showing the influence of environmental variables